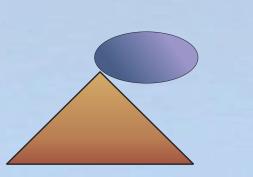
T21A-4558

# Mongolian Hangay Uplift Recorded in Vesicular Basalts



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Sample ID

8.43

8.43

0.52

19.72

DS13-6

DS13-6

DS13-8

DS13-9

DS13-12

DS13-14

DS13-17

DS13-11

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# Abstract

Epeirogenic histories of highland areas have confounded geophysicists for decades, as there are few records paleoelevation in eroding highlands. However, preserved basaltic lava flows record paleoelevation in the size distributions of vesicles at the tops and bottoms of flow units. Although the bubbles have identical mass distributions at top and base, they are subject to different total pressures due to differences in overburden. At the top of the flow, only the overlying atmospheric exerts pressure, Patm, whereas at the base, the lava itself exerts pressure as a function of its thickness, H. Consequently, two factors control the size o bubbles at the base of the flow: atmospheric pressure and lava weight. Thus, the atmospheric pressure-dependence of vesicle size can be expressed by the ratio of vesicle size modes at the top and bottom of a flow. The atmosphere's paleopressure can thus be determined, because all other variables can be measured, and based on the known relationship of pressure as a function of elevation, a paleoelevation can then be calculated. Knowing the elevation at which the rock formed, its age, and its present elevation, the amount of uplift or subsidence can be of the locality. As with any proxy for paleoelevation, there are numerous limitations and error sources. These include the

possibility that the lava was inflated of deflated after solidification of vesicular top and bottom, errors in flow thickness measurement, variations in atmospheric pressure during solidification, the presence or absence of large vesicles in sample cores, and many other minor sources. The total error bounds of the method are estimated to be ±400 m, which is sufficient only for major epeirogenic trends, such as that seen previously on the Colorado Plateau, and now in Mongolia.

As part of a broader collaborative project, we collected samples from several flows from throughout the Hangay Plateau. Results indicate that Hangay experienced uplift of over 1 km in the last 10 Ma at a rate of about 140 m/Ma. A flow sampled from the neighboring Gobi Desert indicates a paleoelevation of only a few hundred m, suggesting that the Hangay Plateau uplifted independently from the regions to the south, which have not experienced the same tectonics as Hangay over the past 20 m.y. The uplift history of Hangay, in addition to the composition of the lavas, geomorphology of the region, drainage pattern history, and other proxies, bears on possible mechanisms for uplift of this part of central Asia, now being explored.

### Introduction

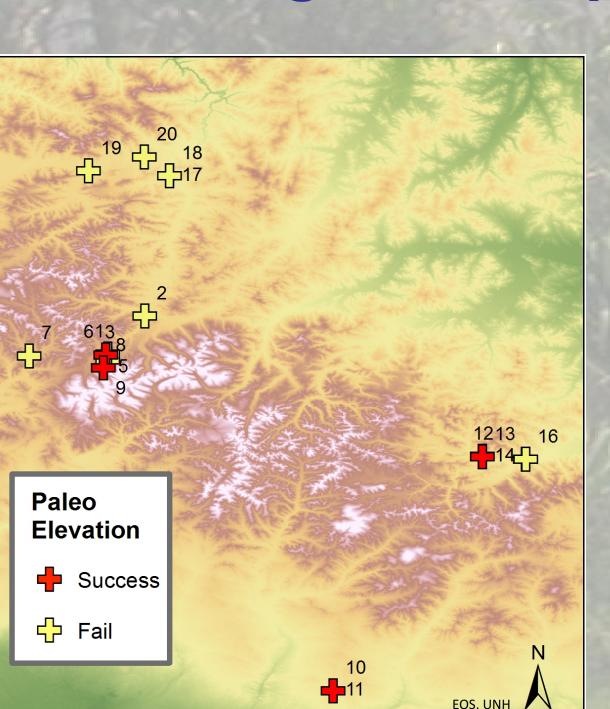
driving mechanisms for uplift remain unclear in many cases. Explanations proposed for their origin have been as diverse as their tectonic settings, and include dynamic forms of lithospheric "drips" involving foundering or delamination of mantle or crustal

The interior of the Hangay Dome in Mongolia locally reaches elevations over 4000 m and contains high-elevation low-relief topography interpreted as an uplifted and well-preserved paleo-erosional surface of presumed Late Cretaceous age, cut into as a measure of atmospheric pressure at the time of eruption, and thus as a

## Sample Dating

ating method. Data was reduced using ArArCALC.

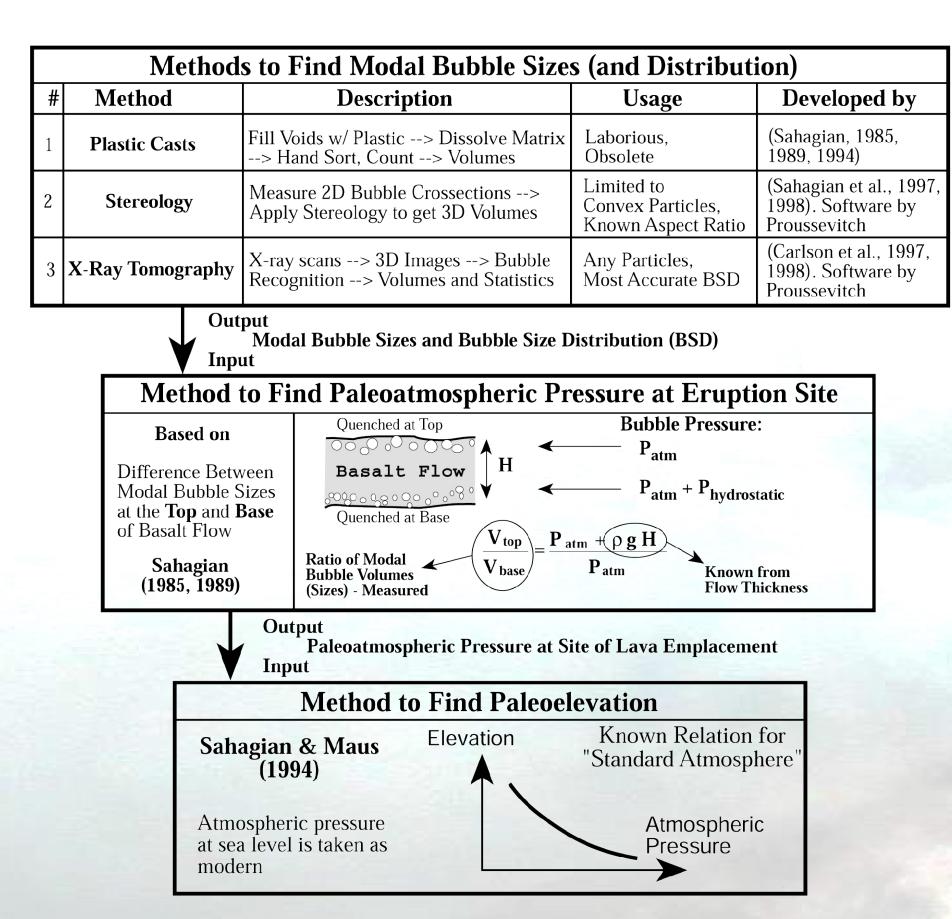
## 2013 Mongolia Samples for Vesicular Basalt Paleoaltimetry



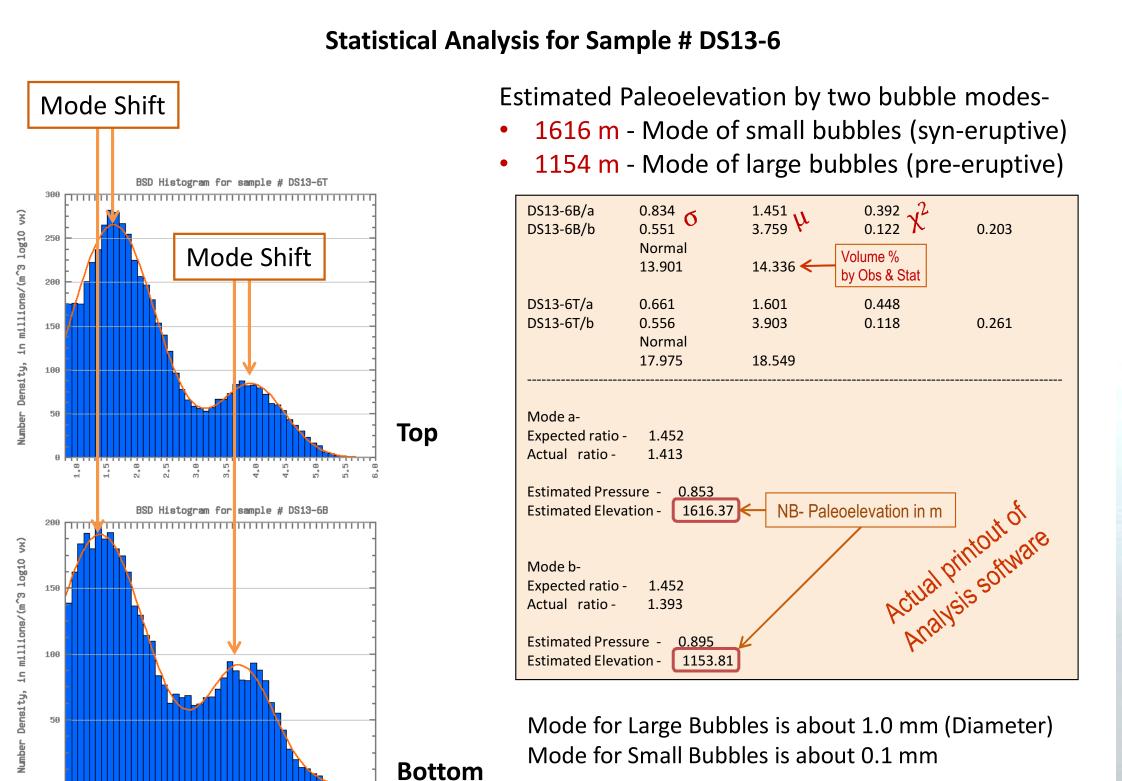
| Date      | Sample ID     | Priority<br>Rating | rating explanation   | Location description  | AGE  | Dating sample | Thickness    | Longitude     | Latitude       | Elevation, r |
|-----------|---------------|--------------------|--|---|------|---------------|--------------|---------------|----------------|--------------|
| 6/21/2013 | DS13-1 (B/T)  | 2                  | Some big vesicles in upper zone.                             | Western slope of Keck camp.   |      |               | 11' 10"      | N 47° 16.457' | E 100° 01.812' | 2404         |
| 6/23/2013 | DS13-2 (B/T)  | 2                  | thick flow   | Western slope of Chulut river   | 100  | Yes           | 22' 3"       | N 47° 26.886' | E 100° 12.051' | 2119         |
| 6/23/2013 | DS13-3 (B/T)  | 2                  | lensy  | Upper lensed flow at the top of thick lava flows above keck camp. Lens width = 20 m. Pipe vesicls at the bottom transitional to 2-3 cm diameter bubbly "pipes" at the middle of flow. |      |               | 7' 3"        | N 47° 16.144' | E 100° 01.254' | 2443         |
| 6/24/2013 | DS13-4 (B/T)  | 4                  | messed up bottom. Sheared upper vesicles                     | Messed up caterpillar bottom. Top has sheared vesicles.   |      | yes           | 12' 4"       | N 47° 14.700' | E 100° 01.897' | 2524         |
| 6/25/2013 | DS13-5 (B/T)  | 2                  | thin flow  | 5 and 6 are two thin flows 1.5 km west of keck camp, 20 m west of DS13-3.   |      |               | 31"          | N 47° 16.144' | E 100° 01.254' | 2443         |
| 6/25/2013 | DS13-6 (B/T)  | 1                  | pretty good  | 5 and 6 are two thin flows 1.5 km west of keck camp, 20 m west of DS13-3.   |      |               | 4' 2"        | N 47° 16.144' | E 100° 01.254' | 2443         |
| 6/25/2013 | DS13-7 (B/T)  | 3                  | big upper vesicles   | Top section of a stack of flows. Big vesicles? Odd - maybe deflated?  | 78   |               | 7' 7"        | N 47° 15.908' | E 99° 40.081'  | 2500         |
| 6/26/2013 | DS13-8 (B/T)  | 2                  | big upper vesicles in core                                   | Mouth of tributary on east side of Baytruk river (Campsite #4). Stack of at least 5 lava flows.   | 例    | yes           | 7' 7"        | N 47° 12.594' | E 100° 00.630' | 2240         |
| 5/26/2013 | DS13-9 (B/T)  | 1                  | pretty good  | next in stack above DS13-8  | 0.00 | yes           | 7' 5"        | N 47° 12.594' | E 100° 00.630' | 2240         |
| 6/27/2013 | DS13-10 (B/T) | 3                  | big upper vesicles   | Gobi- stack of several flows total 50 thick. Bottom two flows are massive (10 m thick) with pillows and massive cores. We sampled next two flows.                                     | il   | yes           | 7' 4"        | N 45° 43.006' | E 101° 04.311' | 1933         |
| 6/27/2013 | DS13-11 (B/T) | 1                  | pretty good  | next in stack above DS13-10   | 475  | yes           | 8' 9"        | N 45° 43.006' | E 101° 04.311' | 1933         |
| 7/2/2013  | DS13-12 (B/T) | 1                  | pretty good  | Stack of multiple flows on south side of Orkon river valley, above rubble slope. 12 is below 13 and 14.   | , 3  | yes           | 10' 8"       | N 46° 47.994' | E 101° 45.643' | 1951         |
| 7/2/2013  | DS13-13 (B/T) | 2                  | thin flow  | middle flow of stack  |      | 1000          | 1' 4"        | N 46° 47.994' | E 101° 45.643' | 1951         |
| 7/2/2013  | DS13-14 (B/T) | 1                  | pretty good  | top flow of stack   |      | yes           | 10' 10"      | N 46° 47.994' | E 101° 45.643' | 1951         |
| 7/2/2013  | DS13-15 (B/T) | 4                  | thick flow, not well measured thickness                      | Orkon waterfall, near base, below very thick flow.  |      | yes           | 20' (approx) | N 46° 47.273' | E 101° 57.616' | 1793         |
| 7/2/2013  | DS13-16 (B/T) | 1                  | pretty good  | Thin flow, near top of cliff, on east side of canyon, at access path  |      | yes           | 6' 8"        | N 46° 47.323' | E 101° 57.676' | 1804         |
| 7/3/2013  | DS13-17 (B/T) | 2                  | lenses, but far from source-<br>good                         | High way bridge at upper end of canyon toward Tariat.<br>Two lenses about 30-50 m long on south side of river. 17<br>below 18.  | 199  | yes           | 5 11"        | N 48° 05.770' | E 100° 18.952' | 1833         |
| 7/3/2013  | DS13-18 (B/T) | 2                  | lenses, but far from source.<br>Heterogeneous upper vesicles | just above DS13/17  |      |               | 12' 1"       | N 48° 05.770' | E 100° 18.952' | 1833         |
| 7/4/2013  | DS13-19 (B/T) | 3                  | not quite top  | East cliff along river of Ger Camp "diamond camp". Stack of flows. Most were unusable. Top sample from DS13/19 was about 20 cm down from pahoehoe top of flow.                        | H    | yes           | 10' 4"       | N 48° 07.133' | E 99° 56.500'  | 2052         |
| 7/4/2013  | DS13-20 (B/T) | 1                  | pretty good  | North cliff of Tariat river canyon. Second flow from top.   |      | yes           | 12' 0"       | N 48° 11.056' | E 100° 12.008' | 1935         |

# Sample Analysis

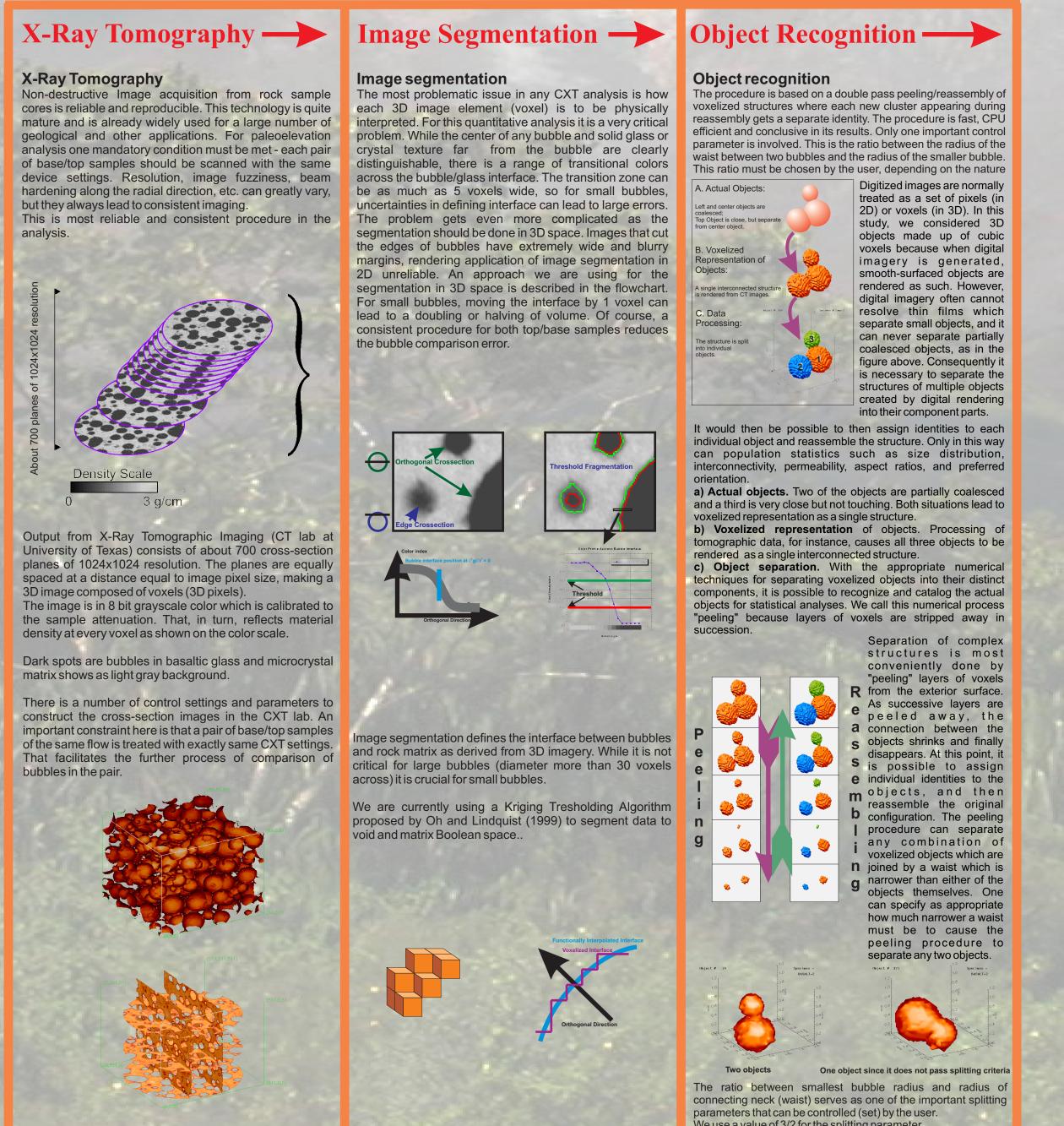
# Principles of Vesicular Basalt Paleoaltimetry

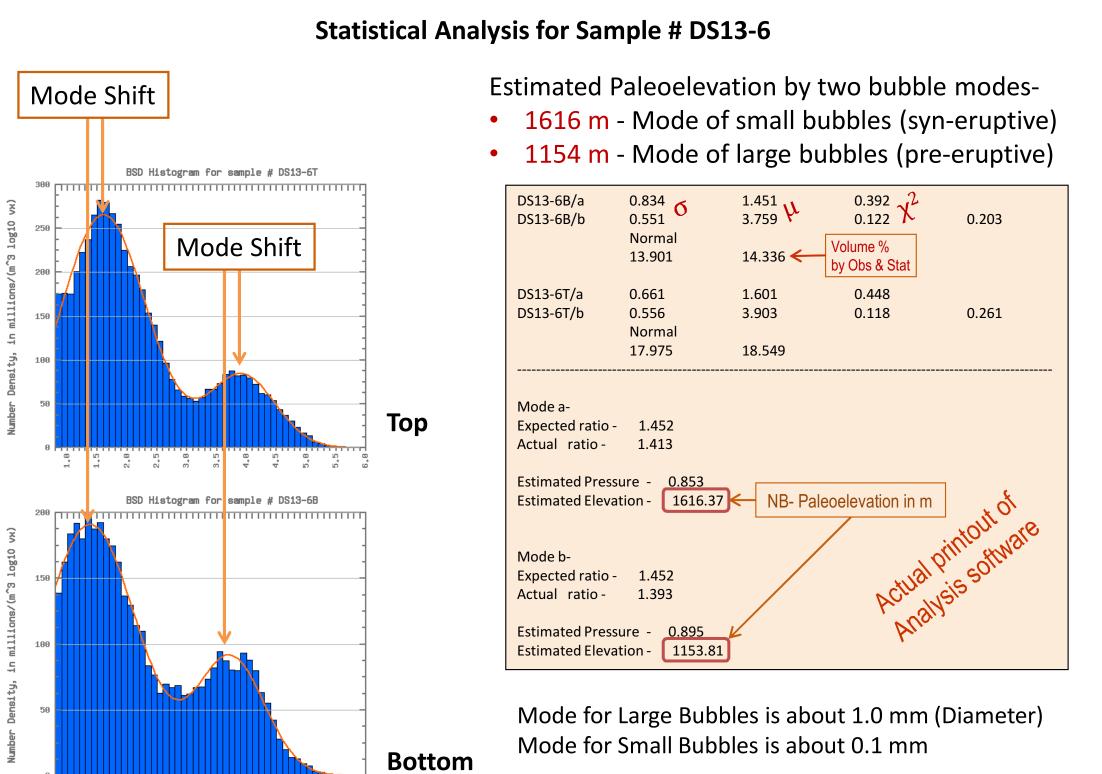


### Sample Processing to Determine Paleoelevation



# Sample Processing to Characterize Bubble Population





2013 Mongolia Samples for Vesicular Basalt Paleoaltimetry (DS13-6)

### Conclusions for acquisition

Bubble size distributions could be obtained from 2D cross-sections or 3D computed X-Ray tomography (CXT) imagery. These techniques to not directly provide mage (2D or 3D) processing for individual bubbles involves segmentation and object Segmentation is the process used to identify each pixel or voxel as void or solid

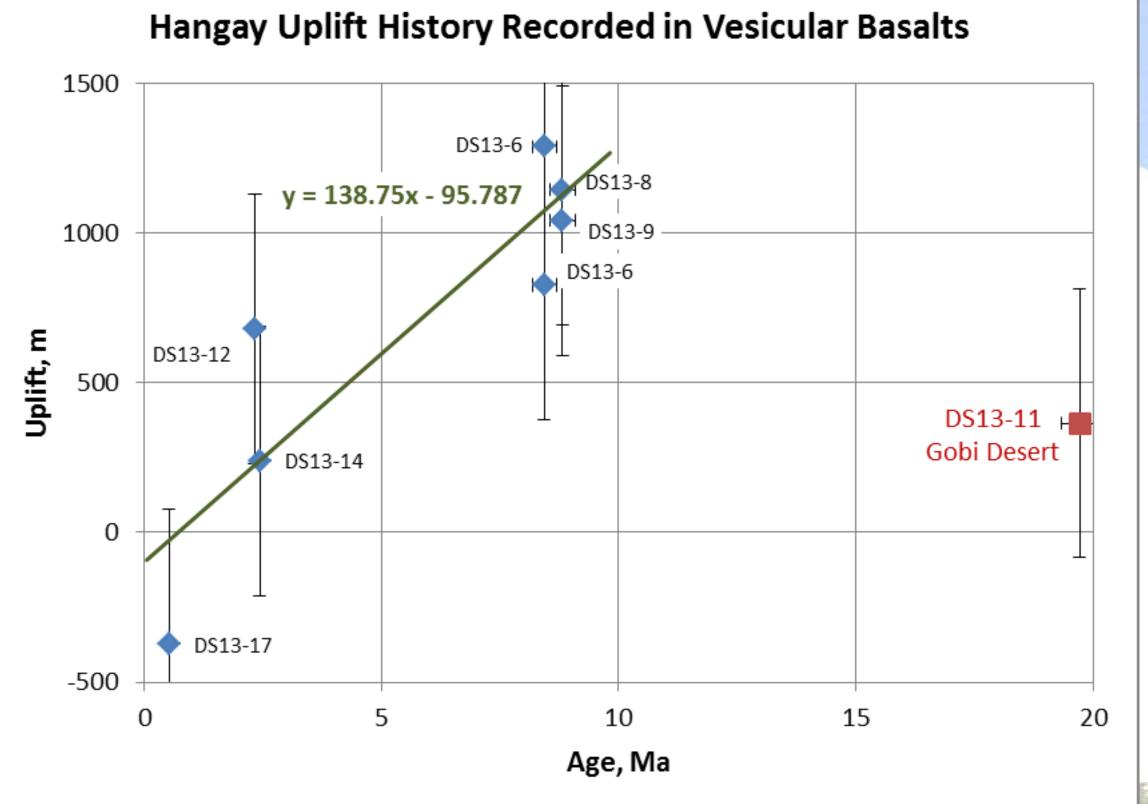
art of the sample. Simple color thresholding proves unsatisfactory, and special nethods of color histogram and Kriging analysis often provide better results. Object recognition is always required to identify bubbles in the segmenter mages since they are often interconnected into clusters as continuous void space Peeling, medial axis analysis, and sphere fitting techniques are currently available to

and subsequent object recognition of 3D CXT imagery directly rovide "true" data for statistical analysis in volumetric (3D) units. Same source data obtained from 2D images represent bubble cross-sections that rarely go through ections needs to be converted to actual bubble sizes in 3D. Stereology is the field of statistical sciences that is directed to perform the kind of conversion that had been nce only applicable to spherical particles. In our previous work we have developed a ormulation for stereological analysis of some basic non-spherical shapes, but nononotonic cross-section frequency functions for these shapes made them virtually larger than their source. So, small observational errors stemming from the inaccurate nature of integer counting for the bins were magnified and made the results

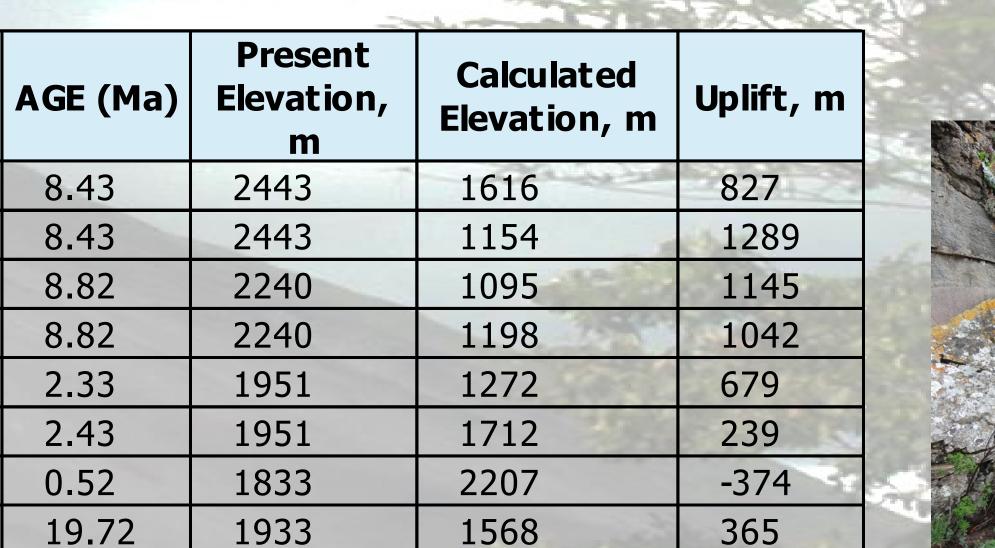
In this work we have developed a practical approach to bypass the problem with nonthis with maximum probability. Testing of this method on the samples with mono- and showed good match. Mono-modal distribution for moderately deformed bubbles rielded almost perfect match while bimodal distribution for highly deformed bubbles is satisfactory but not yet as good to substitute acquisition in 3D for precision

# Results

### Paleoelevation and Uplift



The paleoelevations resulting form analysis of basalt samples collected from throughout the Hangay Plateau indicate about 1 km of uplift in the last 10 million years. While the number of samples is insufficient to resolve acelerations or decelerations of epeirogenic activity, a linear approximation would provide an uplift rate of about 140 m/Ma since the oldest basalts were deposited. The small negative value for DS13-17, a young (0.5 Ma) flow, cannot be distinguished from zero within its error bars, but if the paleoelevation was indeed higher than present, would mean that after a period of uplift, the Hangay recently began to subside, and that the total uplift prior to this was more than measured form the other, uplifted samples. A single, older, sample from the Gobi Desert to the south of Hangay indicates a paleoelevation close to present elevation suggesting that this part of the Gobi did not partake in the uplift experienced by the Hangay.







# Discussion

The overarching goal of this project is to determine the drivers and mechanisms for plateau formation in central Mongolia. The enigmatic nature and position of the Hangay Plateau suggests that heretofore unidentified processes are operating on this part of the Asian lithosphere.

The constraints provided by the vesicular basalts sampled in this part of the project suggest that substantial uplift occurred in the last 10 Ma, rather that being limited to the initial period of India-Asia convergence and mountain-building. As such, more localized mechanisms involving mantle-lithosphere interactions are likely to be playing a role. As results from sister studies involving seismics, geochemistry, geomorphology, isotope systematics, and genetic drift are concatenated, closer constraints on the mechanism of plateau uplift operating in Mongolia should emerge.





Acknowledgments

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